

GEOLOGIC MAP OF THE HUACHUCA CITY 7 1/2' QUADRANGLE, COCHISE COUNTY, ARIZONA

Philip A. Pearthree and Ann Youberg

Arizona Geological Survey Digital Geologic Map 36 (DGM-36), version 2.0

April 2009

1:24,000 scale

Citation for this map:

Pearthree, P.A., and Youberg, A., 2009, Geologic Map of the Huachuca City 7 1/2' Quadrangle, Cochise County, Arizona: Arizona Geological Survey Digital Geologic Map 36 (DGM-36), version 2.0, 1 sheet, layout scale 1:24,000.

(also available in Adobe pdf format on CD-ROM)
Not to be reproduced for commercial purposes

Research supported by the U.S. Geological Survey, National Cooperative Geologic Mapping Program, under USGS award #02HQAG0016. The views and conclusions contained in this document are those of the authors and should not be interpreted as necessarily representing the official policies, either expressed or implied, of the U.S. government. The Arizona Department of Water Resources provided funding for map revisions depicted in DGM-36, version 2.0.

Acknowledgements:
Genevieve Pearthree was the cartographer for DGM-36, version 2.0.

Map Unit Descriptions

Other units

- Plowed areas** - Historically or actively plowed fields, irrigated pastures, and other lightly disturbed ground.
- Disturbed ground** - Profoundly disturbed areas, including aggregate excavations and intensively developed trailer parks that completely obscure the underlying geologic units.

Babocomari River alluvium

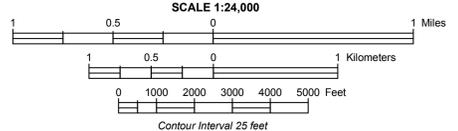
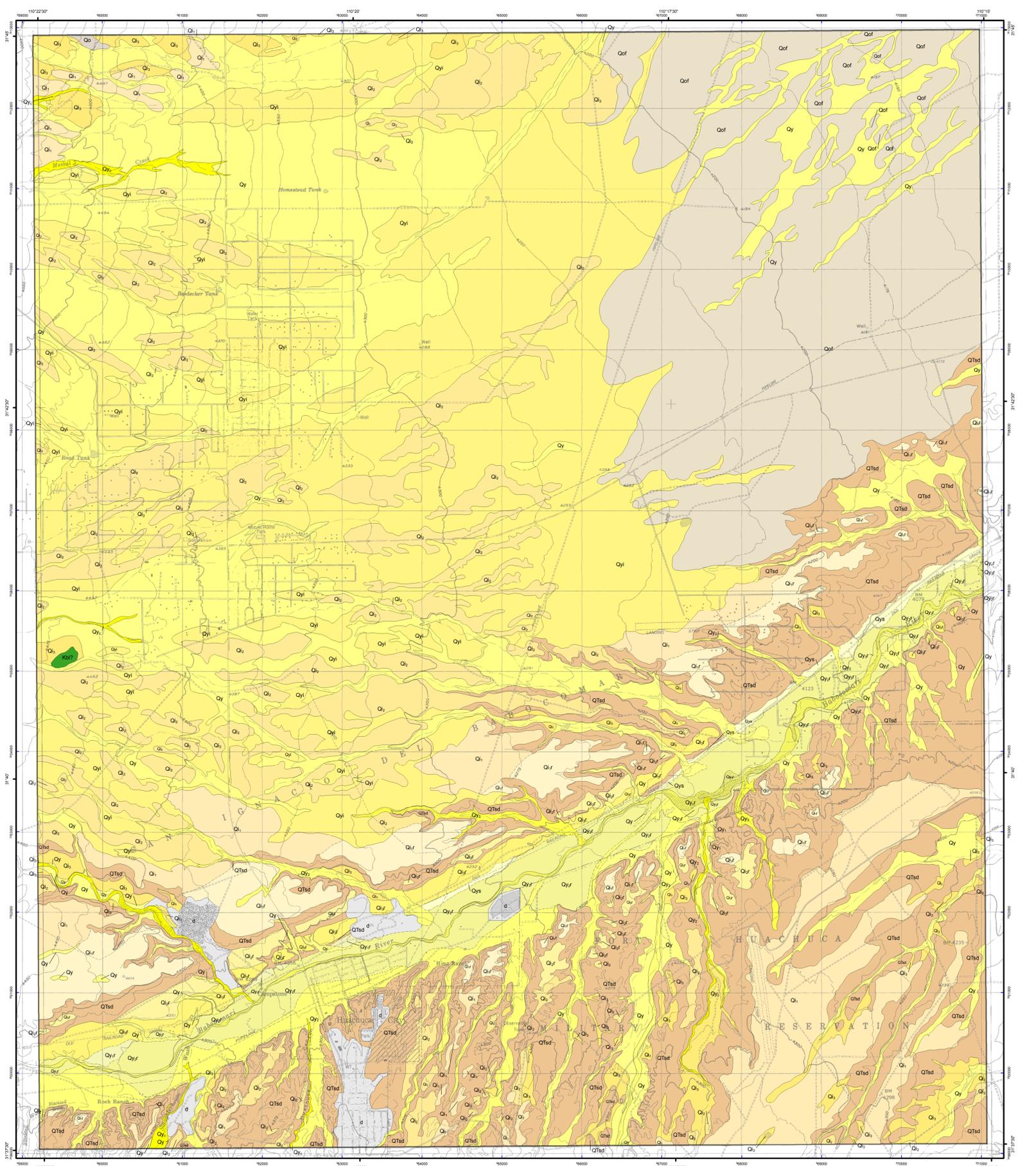
- Active river channel deposits** - Deposits are dominantly unconsolidated, very poorly sorted sandy to cobbly beds exhibiting bar and swale microtopography but can range from fine silty beds to coarse gravelly bars in meandering reaches based on position within the channel. Clasts are typically well-rounded but may be angular to sub-angular. Cyclic deposits are typically unvegetated to lightly vegetated and exhibit no soil development. Cyclic deposits are entrenched from 30 cm to 10 meters or more below adjacent early historical floodplain deposits depending on location, geomorphic relationship, and local channel conditions. Although much of the Babocomari River was a perennial stream historically, some sections are dry or marshy at the channel during much of the year. These deposits are the first to become submerged during flow events and can be subject to deep, high velocity flow and lateral bank erosion. In some areas of the Babocomari River channel deposits are very thin to discontinuous exposing underlying bedrock. Extent of channel deposit and exposed bedrock varies and shifts with significant flooding.
- Historical river terrace deposits** - Terrace deposits that occupy elevations from 1 to 2 meters above Cyclic deposits and are inset below the pre-incision historical floodplain. These surfaces are generally planar but exhibit bar and swale microtopography. Although no soil development is present, dense grasses and small mesquite trees abound. Sediments composing these deposits are poorly sorted silt, sand, pebbles and cobbles. Pebbles and cobbles are well-rounded to sub-angular, ripple marks, and stacked channel deposits are visible in cross-section. Deposition in a low to moderate energy braided stream environment. These deposits are prone to flooding during extreme flow events, and undercutting and rapid erosion of Cyclic surfaces is possible during lower flow events.
- Latest Holocene to historical river terrace deposits** - Deposits associated with the floodplain that existed prior to the early historical entrenchment of the Babocomari River (Hendrix, 1950; Huckleberry, 1998; Wood, 1997). Cyclic deposits are associated with broadly planar surfaces that locally retain the shape of historical river meanders. Cyclic surfaces are up to 7 meters above modern Cyclic deposits and are the most extensive river terraces in the valley. Cyclic sediments were deposited when the San Pedro and Babocomari Rivers were widespread, shallowly-flowing river systems and are dominated by fine grained floodplain deposits. Dense mesquite bosque and tall grass is typically present on these surfaces except where historic plowing or grazing has taken place. These surfaces appear predominantly fine grained at the surface due in part to the input of organic matter and windblown dust deposition but are composed of intertuffing coarse sandy to pebbly braided channel and fine sand to silty river floodplain deposits. Where Cyclic deposits are moderately to deeply incised they are not subject to inundation by river flows, but they may be flood-prone in areas with less channel incision. Cyclic deposits are subject to catastrophic bank failure due to undercutting and lateral erosion during flow events. Distal piedmont fan deposits (Qy2, Qy3f, and Qy3) overlap onto Cyclic deposits although an intertuffing relationship likely exists in the subsurface.
- Late to early Holocene river terrace deposits** - Deposits associated with slightly higher terraces that represent either higher elements of the early historical floodplain or remnants of older Holocene aggradation periods. These fine-grained terrace deposits commonly have been disturbed by plowing or cattle grazing. When undisturbed, Cyclic deposits are densely vegetated by mature mesquite trees (mesquite bosque) and tall grasses. Soil development is moderate and surface color ranges from 10 to 7.5 YR 4/4. Due to the dense vegetation, input of organic matter at the surface often results in a thin (10 cm) organic soil horizon. A light dusting carbonate stage II calcium carbonate accumulation is evident on the undersides of some terrace clasts. Cyclic surfaces are up to 7 m above the active channel in highly incised locales and typically are less than 1.5 m higher than adjacent Cyclic surfaces. Along the Babocomari River Cyclic terraces are found along the valley margins. These terraces typically are covered with fine-grained floodplain deposits, but relict gravel bars and lenses are common.
- Late Pleistocene river terrace deposits** - Unit Q3r consists of remnant river terrace deposits. These terraces are typically 5 to 10 m above the active channel, and in a few areas are clearly separated by a few meters of elevation. Q3r surfaces commonly have loose, open lags of cobbles and gravels; surface clasts exhibit weak rock varnish. Q3r surfaces appear light orange to reddish brown in color, and the surface soil horizon. Q3r soils are moderately developed, with orange to reddish brown sandy loam to clay loam argillic horizons and stage II calcium carbonate accumulation.
- Middle to late Pleistocene river terrace deposits** - Higher intermediate terraces about 10 to 15 m above the Holocene floodplain of the Babocomari River. Terrace surfaces typically are fairly resistant to slumping remnants perched on eroded St. David Formation deposits. Terrace deposits are a mix of river sand, gravel, silt, and clay, but surfaces typically are covered with relict gravel deposits. Soil development is moderate because surfaces are not extensive, consisting primarily of weak clay accumulation and 5 YR reddening and stage II to III calcic horizons.
- Early to middle Pleistocene river terrace deposits** - Unit Q1r consists of higher remnant river terrace deposits. These terraces are typically about 30 m above the active channel. Q1r consists of fine-grained sediment, Q1r1 includes coarse-grained fine-grained sediment. Q1r1 surfaces commonly have loose, open to moderately incised cobbles and gravel surface lags, surface clasts exhibit moderate rock varnish. Q1r1 surfaces appear in orange color aerial photos, reflecting reddening of surface clasts and the surface soil horizon. Soils typically contain reddened, clay argillic horizons, with obvious clay skins and subangular to angular blocky structure. Underlying soil carbonate development is typically stage II, with abundant carbonate through at least 1 m of the soil profile.

Piedmont alluvium and surficial deposits

- Modern stream channel deposits** - Unit Qm consists of deposits in active channels of the larger tributary drainages. Channel deposits are mapped where they are extensive enough to represent at 1:24,000 scale. They were outlined using the 1997 digital orthophotos for the Huachuca City quadrangle. Deposits are composed primarily of sand, pebbles, cobbles, and small boulders. Locally channels are incised as much as 2 to 3 meters below adjacent Holocene terraces (Unit Qy), but in other areas there is no incision. Channels consist of single, relatively large channels and smaller branching channels in areas of channel expansions. Local relief within channels varies from minimal to more than 1 meter between low-flow channels and adjacent gravel bars.
- Latest Holocene alluvium** - Recently active piedmont alluvium located primarily along active drainages including floodplain, low-lying terrace, and intermittent tributary channels. Qy3 deposits are composed of unconsolidated to very weakly consolidated sandy to pebbly deposits and exhibit greater vegetation than Cyclic deposits. These deposits generally exhibit bar and swale microtopography and are susceptible to inundation during moderate to extreme flow conditions when channel flow exceeds capacity. Soil development is generally absent on Qy3 deposits which exhibit pale buff to light brown (10 YR) surface coloration.
- Late Holocene alluvium** - Young deposits in floodplains, low terraces and small channels that are part of the modern drainage system. Along the larger drainages, unit Qy2 sediment is generally poorly to very poorly sorted silt, sand, pebbles, and small cobbles; floodplain and terrace surfaces typically are mantled with sand and finer sediments. On lower alluvial areas and smaller tributaries young deposits consist predominantly of moderately to moderately well-sorted sand and silt, with some pebbles and cobbles in channels. Soils are pale brown in color (10 YR), and soil development is very weak, consisting of slight carbonate accumulation. Channels generally are incised less than 1 m below adjacent terraces, but locally incision may be as much as 2 m. Channel morphologies generally consist of a single-thread channel or multi-threaded channels with low flow channels with gravel bars and low flow channels with fine sand and silt. Channel morphologies are flood prone and may be subject to deep, high velocity flows in large flow events. Potential lateral bank erosion is severe, and flood flows may significantly change channel morphology and flow paths. Local relief varies from fairly smooth channel bottoms to undulating bar-and-sand topography that is characteristic of coarser deposits. Terraces have planar surfaces, but small channels are common.
- Older Holocene alluvium** - Qy1 deposits consist of terraces along tributary drainages and broad, low-relief, undulating fan deposits that exhibit shallow, widespread braided channel networks and higher in the landscape than younger Holocene alluvium. Overall relief between broad fan crests and incised drainages on gently rolling Qy1 deposits typically does not exceed 1.5 meters. Numerous shallow braided channels drain widespread portions of Qy1 surfaces. Qy1 deposits exhibit incipient calcic carbonate accumulation (stage I) and soil development characterized by medium brown (10-7.5 YR) coloration where unincised.
- Fine-grained Holocene alluvium derived from the St. David Formation** - Thin, fine-grained Holocene alluvial deposits mantling valley bottoms. Sediment is primarily clay silt and sand, with occasional gravel, and with occasional fine gravel. Soil development is minimal, where it has developed soil is typically a brown (7.5YR) sandy loam with substantial disseminated carbonate but no visible carbonate accumulation.
- Holocene to late Pleistocene alluvial fan and terrace deposits** - Mixed fine-grained Holocene (Qy) and Pleistocene (Q2 or Q3) alluvium. In some areas, surficial deposits are primarily Holocene, but these deposits are quite thin and some Pleistocene deposits are exposed along channels. In other areas, surficial deposits are primarily Pleistocene but there are Holocene deposits in topographically low areas.
- Holocene alluvium undifferentiated** - Unit Qy consists of young deposits in undivided small channels, floodplains, low terraces, and alluvial fans on the upper piedmont. Deposits vary widely in particle size. Qy deposits on the upper piedmont and along major tributaries are very coarse, including boulders, cobbles, pebbles, sand, and minor silt and clay. Farther downslope, Qy deposits are fine, consisting of sand and silt, with minor fine gravel and clay. Small channels generally are incised less than 1 m below adjacent Qy terraces and fans. Channel morphologies generally consist of a single-thread channel or multi-threaded channels with low flow channels with gravel bars and low flow channels with fine sand and silt. Channel morphologies are flood prone and may be subject to deep, high velocity flows in large flow events. Potential lateral bank erosion is severe, and flood flows may significantly change channel morphology and flow paths. Local relief varies from fairly smooth channel bottoms to undulating bar-and-sand topography that is characteristic of coarser deposits. Terraces have planar surfaces, but small channels are common.
- Late Pleistocene alluvial fan and terrace deposits** - Unit Q3 consists of deposits associated with moderately dissected terraces and small relict alluvial fans found on the upper, middle and lower piedmont. It also includes some deposits on topsets of dissected ridges in the Babocomari Valley. Moderately well developed, moderately incised tributary drainage networks are typical on Q3 surfaces. Active channels are less than 5 meters below Q3 surfaces. Q3 fans and terraces are lower in elevation than adjacent Q2 or older surfaces, but elevation differences are minimal in some places. Q3 deposits generally consist of pebbles, cobbles, and finer-grained sediment. Q3 surfaces commonly have loose, open lags of pebbles and cobbles; surface clasts exhibit weak rock varnish. Q3 surfaces are characterized by light reddening of surface clasts and the surface soil horizon. Q3 soils are moderately strongly developed, with orange to reddish brown clay loam to light clay argillic horizons and stage II calcium carbonate accumulation.
- Middle to late Pleistocene alluvial fan and terrace deposits** - Unit Q2 consists of moderately dissected relict alluvial fans and terraces with strong soil development found throughout the map area. Q2 surfaces are drained by well-developed, moderately to deeply incised tributary channel networks; channels are typically several meters to as much as 10 m below adjacent Q2 surfaces. Q2 deposits typically consist of sand, pebbles and cobbles. Q2 surfaces are characterized by scattered cobble to boulder lags with moderate to strong varnish. Well-preserved, planar Q2 surfaces are smooth with scattered pebble and cobble lags; surface color is reddish brown and rock varnish on surface clasts is typically orange or dark brown. More eroded, rounded Q2 deposits are less clay-rich and have some carbonate litter on the surface. Well-preserved Q2 surfaces have a distinctive bright red color on color aerial photos, reflecting reddening of the surface soil and surface clasts. Soils typically contain red clay argillic horizons, with obvious clay skins and sub-angular to angular blocky structure. Underlying soil carbonate development is typically stage II to III, with abundant carbonate to depths of at least 1.5 m in the soil profile, but indurated petrocalcic horizons were not observed.
- Early to middle Pleistocene alluvial fan and terrace deposits** - Unit Q1 consists of deposits associated with deeply dissected relict alluvial fans with variable soil development. Q1 surfaces are drained by well-developed, deeply incised tributary channel networks. Q1 surfaces are typically 5 to 20 meters above adjacent active channels. Q1 deposits are not well-exposed, but probably consist primarily of pebbles, cobbles, sand, and silt and clay. Where Q1 surfaces are fairly well preserved, they are smooth with pebble and cobble lags; rock varnish on surface clasts is typically orange to red. Soil carbonate development is variable, but typically is stage III. Where surfaces are well preserved, soils typically contain dark red, clay loam argillic horizons, with obvious clay skins and sub-angular blocky structure. More eroded Q1 surfaces are characterized by loose cobble lags with moderate to strong varnish, ridge-and-valley topography, and carbonate litter on the side slopes. On aerial photos, ridge crests on Q1 surfaces typically are dark reddish brown, reflecting reddening of the surface soil and surface clasts, and eroded slopes are gray to white.

Bedrock units

- Limestone of the Cretaceous Bisbee Group (7)** - A sequence of thin- to medium-bedded, gently northwest-dipping gray limestone, locally finely laminated. Some chert concretions and quartz veins are locally abundant. Dark gray to black Neohy limestone outcrops mapped in the Mustang Mountains include Colina Limestone and Horquilla Limestone of the Naco Group and Bisbee limestone (Hayes and Raup, 1968), characteristics of limited outcrop seem to fit Bisbee limestone best.

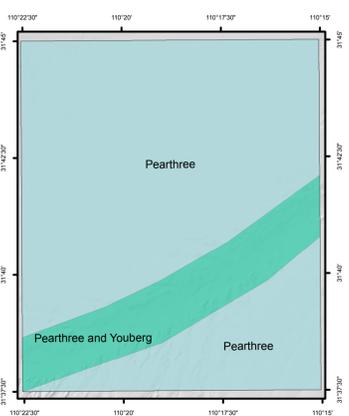


Arizona Geological Survey
416 W. Congress Street, Suite 100
Tucson, AZ 85711
(202) 770-3500
www.azgs.az.gov

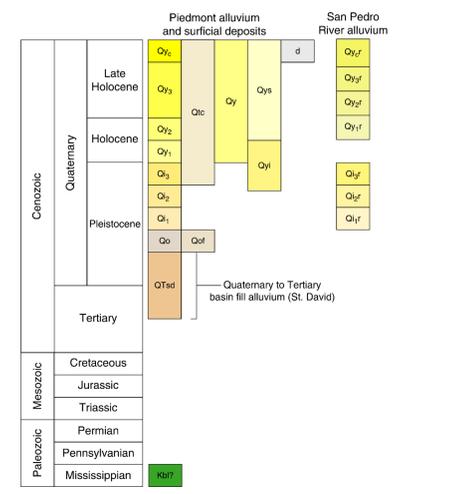


Topographic base from USGS 1:24,000 scale quadrangle series. North American Datum of 1983 (NAD83). Projection and 1,000-meter grid: Universal Transverse Mercator, zone 12.

Mapping Responsibility



Unit Correlation



Map Symbols

- Contacts**
- accurate contact
- approximate contact
- concealed contact

Cochise County

Mapped Area Shown in Blue

Location Index Map

Quadrangle Location Shown in Blue



Adjoining 7.5' Quadrangles

